

Landslide hazard zonation assessment using GIS analysis at the coastal area of Safi (Morocco)

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Abstract: Landslide hazard is one of the major environmental hazards in the coastal area. For helping the planners in selection of suitable locations to implement development projects, a landslide hazard zonation map has been produced for the coastal area of Safi (Morocco) as part of coastal Meseta. For this purpose, after preparation of a landslide inventory of the study area, some major parameters were examined for integrated analysis of landslide hazard in the region. The analyses of parameters were done by geo-referencing and lateral model making, satellite imaging of the study area, and spatial analyses by using geographical information system (GIS). The produced factor maps were weighted with analytic hierarchy process (AHP) method and then classified. The study area was classified into four classes of relative landslide hazards: negligible, low, moderate, and high. The final produced map for landslide hazard zonation in coastal area revealed that: the parameters of slope, geologic formation and fracturation have strong correlation and predict 75% of existing instabilities.

Keywords: landslide hazard zonation map, geographical information system (GIS), analytic hierarchy process (AHP), Coastal area, Safi, Morocco

1. Introduction

Landslide are belonging to the most damaging natural hazards. Preventing damage requires a reliable predictive inventory of the dynamic processes and the implementation of risk maps, necessary as base for any development. It is essential to define the factors and mechanisms that cause the movements in order to contribute to a better approach for the detection of the phenomena of instability and to risk assessment.

The assessments of landslide hazard and susceptibility are the relative spatial probability of a new landslide occurring in the future (Remondo et al., 2003), and their assessment in a given area should normally be based on the analysis of instability factors. Using geographical information system (GIS) analysis provides a powerful tool to model the landslide hazards for their spatial analysis and prediction (Dai et al., 2002; Cevik and Topal, 2003; Ayalew and Yamagishi, 2005; and Fall et al., 2006). Thus, it is necessary to look at the conditions under which landslide have occurred in the past, and to use the critical combinations of preparatory factors for delineating the possible occurrence of further landslides. The most common natural triggering factors are intense rainstorms, prolonged periods of wet weather, or seismic activities. Stronger earthquakes could be a triggering factor as well as long-term, neotectonic movements.

Several algorithms and models have been proposed for generating the Landslide Susceptibility that mainly include Analytical Hierarchy Process (AHP) (Saaty, 1980; Saaty and Vargas, 2001), logistic regression (Carrara et al., 1983), fuzzy-logic (Gee et al., 1991), artificial neural

network analysis (Ermini et al. 2005; Regmi et al. 2010; Ruff and Rohn 2008, Canani et al., 2008). modeling approaches (Perriello Zampelli et al., 2012), Fuzzy Analytical Hierarchy Process (FAHP) (Shadman Roodposhti et al., 2010), Geographically weighted principal component analysis (Faraji Sabokbar et al., 2014), GIS matrix method (GMM) (Irigaray 1995; Irigaray et al. 1999, 2007; Clerici et al. 2002; Jiménez-Perálvarez et al. 2009; Jiménez-Perálvarez 2012 etc, that most of which are related to the weight of landslide factors. Above studies demonstrate that many techniques have been used for landslide susceptibility mapping and have achieved excellent results. One of these methods is the AHP that was used.

AHP gives a proven, effective means to deal with complex decision making and can assist with identifying and weighting selection criteria, analyzing the data collected for the criteria and expediting the decision-making process. AHP has been shown to be the best method of eliciting and using multi criteria preference relationships in a range of applications. The AHP is based on a matrix of pair wise comparisons between criteria, and it can be used to evaluate the relative performance of decision alternatives (for example products and services) with respect to the relevant criteria. The AHP was seen to be a suitable tool for the purpose here, as it is a method that is particularly suited to decisions made with limited information (Saaty et al. 2001).

AHP allow some small inconsistency in judgment because human is not always consistent.

In this paper, AHP (Saaty, 1980; Saaty and Vargas, 2001) was adopted to establish landslide susceptibility maps.

The AHP method is used to determine the weights of various themes for identifying the landslide area based on weights assignment and normalization with respect to the relative contribution of the different themes to landslide occurrence for classified into four categories low, medium, high and very high potential zone identify through integration of various thematic maps with GIS techniques (Jhariya et al , 2016). The weights of each factor were computed statistically using multi-influencing factor technique followed by heuristic approaches driven method for assigning ranks to each sub-classes of factor maps at resulted for AHP model. In this context, multi-criteria decision making (MCDM) is rather a simple, effective and reliable technique for AHP process for the identification of landslide hazards areas.

A multi parametric data set comprising remote sensing data and conventional maps at the caterious wise all five different thematic layers were integrated with weighted overlay in GIS. For analysis result to different of thematic layers, namely, geomorphology, geology, soil type, slope, and land use/land cover, are prepared using satellite imageries, topographic maps, and secondary data set, and integrated with weighted overlay in GIS to generated landslide zonation map.

2. Materials and Methods

2.1 Study area

The study area is located in the coastal area of Moroccan atlantic ocean (Figure1). It covers an area of about 4110 Km² between longitudes 09°20'W and 08°51'E, and latitudes 32°24'N and 32°68'N. The geomorphology of the studied area is characterized by the abundance of consolidated sand- stone of a Plio-Quaternary age, forming the dune reliefs parallel to the ocean, with a SSW-NNE orientation

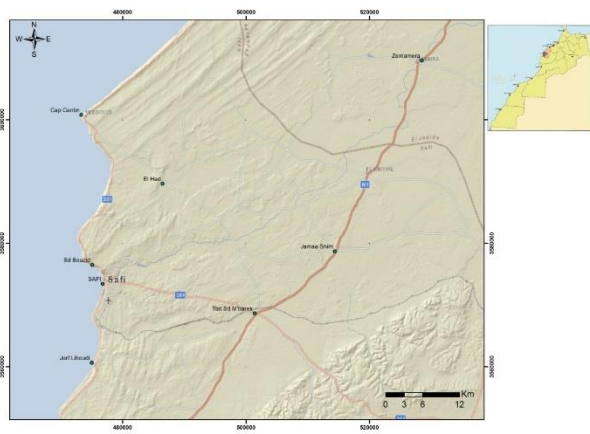


Figure 1: Location of study area

The atlantic moroccan coast of safi is characterized by longest (60 Km) and high live cliffs. They are developed in the mesozoic layers, made up of alternate beds of clay, marls and limestones wich give different sections vertical in parts only. The processes of evolution of the cliffs as a consequence of landslides are now active and dangerous.

The altitude varies from 500 to 50 m, and it decreases from the southeast to the northwest. The study area has a semi-arid climate with mean annual precipitation of 325 mm and mean annual temperature of 22°C. The rainiest month is November and the driest month is August.

This area belongs to the mesetian domain characterized by Meso-Cenozoic tabular formations deposited on the Paleozoic basement (Roch 1950; Gigout 1951; Michard 1976; Saaidi 1988; Witam 1988; Ouadia 1998).

The encountered geological formations are the Evaporitic Complex and Yellow Dolomite (Upper Jurassic), Lower Limstone (Late Berriasian), Brown Clay (basal Valanginian), Dridrat Limestone (Hauterivian), The "red sandy clay beds" of the Upper Hauterivian and clastic and Shelly limestone (Plio-Quaternary); (Figure 2). These geological formations are the most sensitive to landslide. However, the formation most susceptible to sliding is Brown Clay localized in the north coastal cliffs.

The tectonic structure is essentially tertiary in age, characterized by two main directions of folds: NS and SW-NE to EW, inherited respectively from Hercynian and Alpine orogeny (Ferré and Ruhard 1975). These tectonic directions guide and control the development of landslide structures

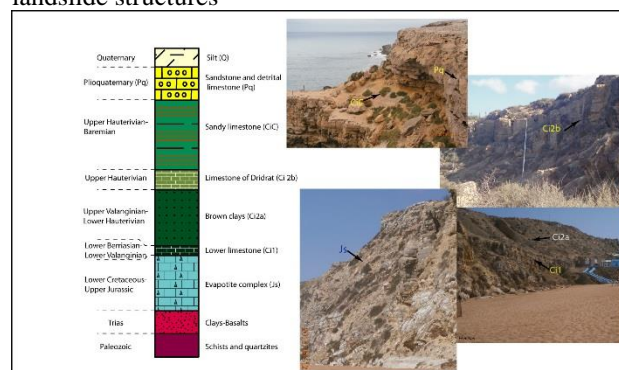


Figure 2: Geological setting of Safi region

2.2 Data

The occurrence of landslides is a function of direct and indirect natural and human factors.

In this study, a spatial database considering landslide-related factors (Figure 3) based on extensive field work and previous inventory maps, such as topography, geology, and land cover, was used. Land cover was detected from satellite images such as TM images. The landslide map is in the form of point coverage, the topographic map in the form of line and point coverage at a scale of 1:50,000, the geological map in the form of polygon coverage at a scale of 1:50,000. Eight factors, extracted from the constructed spatial database, were considered when calculating the probability.

The slope, aspect, were obtained from the Digital Elevation Model (DEM). The distance from stream and road was calculated using the topographic map. A lithology and fault map of the study area is digitized from the existing geology map at the scale of 1:200,000. A fracturation map was extracted from Landsat TM image using filter processing. Finally, land cover data was classified from a LANDSAT TM image using the supervised (maximum likelihood) classification method

The precipitation is also one of major parameters because Most of the landslides occur after heavy rainfall; so that water infiltrates rapidly upon heavy rainfall and increases the degree of saturation and potential of landslide occurrence [Pourghasemi et al., 2009].

The main factors affect the susceptibility mappings are: landslide inventory, precipitation, lithology-geology, aspect, slope, land cover, distance to stream, and distance to road (Figure 3).

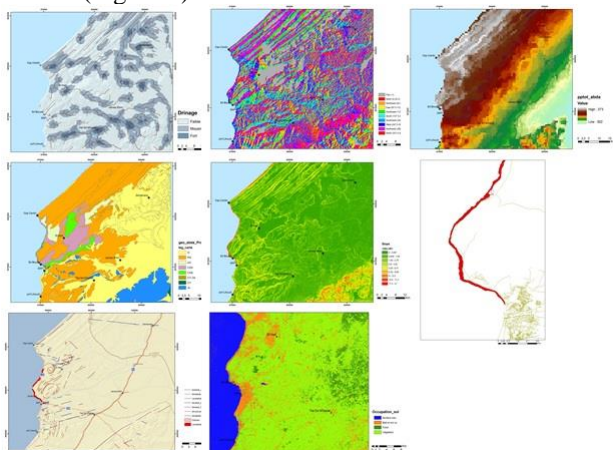


Figure 3 : Example of input Data

2.2.1. Landslide inventory

It is very important to determine the location and area of the landslide correctly when preparing the landslide susceptibility maps. Landslide inventory is an essential part and basic information for any landslide zoning such as susceptibility, risk and hazard zonings. It involves the location, classification, volume, travel distance, state of activity and date of occurrence of land sliding in an area (Fell et al. 2008a) There are different methods to identify landslides. They include aerial photos, satellite images, literature survey for historical landslide records and field observations (Figure 4).



Figure 4; Example of Landslide of the study area

The landslide inventory was determined from field records and visual interpretation of aerial photographs and satellite images. In this study, the field record has been used to determine 170 landslides.

The majority of landslides identified in this region have been occurred in cliffs banks and roads. The landslides in cliffs are often due to loss of land cover, steep slope topographic, and erosions. Road construction has been caused of many landslides in this region as well.

2.2.2. Lithology-geology

The main source of data related to the geomorphology of an area of land is determined by the lithology properties of that land (Dai et al. 2001). The landslide phenomenon, a part of the geologic studies and research, is related to the lithology and weathering properties of the material of the land (Yalcin, 2008). In the most of recent studies, such as references (Ayalew et Yamagishi, 2005) and (Yalcin, 2008), this parameter has been considered as the most important factor in landslide susceptibility mapping.

The most important formations in this region having high potential for occurring landslide are Brown Clays (Ci2a). These formations are commonly composed of marl and clay rocks.

Marl has high cohesion and is very sensitive in neighborhood of water. In these formations, landslide occurs when water infiltrates in internal layers and external layers are exposed to weathering. In this region, the rocks are exposed to physical and chemical factors. in order to consider the effect of lithology and geology, this region has been divided into 7 parts.

2.2.3. Slope

The slope is one of the main parameters in the slope stability analysis. The slope angle directly affects landslide; thus it is used in preparing landslide susceptibility maps (Pant, M., 1992, Lee et al., 2004a, Lee, 2005, Saha et al., 2002).

In some of the recent studies, such as by Yao et al. 2008 and Nandi et Shakoore 2009, this parameter has been considered as the most important factors in landslide susceptibility mapping. For preparing landslide susceptibility map, the slope map was divided into nine slope categories (See Table. 2). According to the landslide inventory map, most landslides had occurred in 17-31 degree of slope ranges.

2.2.4. Aspect

Despite the fact that some authors mentioned that aspect has no significant influence on land sliding, several researchers have reported a relationship between slope orientation and landslide occurrence (DeGraff et al., 1980, Marston et al. 1998). Aspect-related parameters such as exposure to sunlight, drying winds, rainfall (wetness or degree of saturation) and discontinuities may control the occurrence of landslides. In this study slope aspect is divided every 45°. Thus there are 8 intervals

2.2.5. Land cover

Land use/land cover are also key factors responsible for landslide occurrences.

Generally, land cover has effect on strength of slope materials against sliding and control of water content of

slope. In addition, plant roots reinforce the slope and normally are considered as reinforcements. Land cover absorbs the water of soil and decreases the potential of landslide. This is an important issue in marl soils. In studies performed by Komac 2006 and Leventhal et Pant, 1992, this parameter has been considered as one of the most important factors in preparing landslide susceptibility maps. Land cover is determined from satellite images and field investigations. In this study, land cover map is divided into four land cover categories.

2.2.6. Precipitation

Rainfalls produce sudden floods which cause shallow landslides. Most of the landslides occur after the heavy; so that water infiltrates rapidly upon heavy rainfall and increases the degree of saturation and potential of landslide occurrence (Pourghasemi et al., 2009).

2.2.7. Fracturation

Fracturation are the structural features, which describe a zone of weakness with relative movement, along which landslide susceptibility is higher (Pourghasemi et al, 2009). It has generally been observed that the probability of landslide occurrence increases near of the fracturation, which not only affect the surface material structures but also make contribution to terrain permeability causing slope instability (Iovine et al, 2008)

2.2.8. Distance from road

The distance to road parameter reflects human activities. In other words, landslides may occur on the slopes intersected by roads (Nielsen et al, 1979). A road or power line constructed beside slopes causes a decrease in the load on both the topography and the heel of slope. According to recent studies, cutting slopes for road construction and frequency vibrations caused by cars would induce landslides (Mittal et al., 2008). During field studies, some landslides were observed, which certify the effect of instability caused by road constructions (Figure. 5).



Figure 5: Road near Safi prone to landslide

2.2.9. Distance from streams

Generally, potential of landslides increases by decrease in distance to streams, because streams may adversely affect stability by eroding the slopes or by saturating the lower part of material, resulting in water level increases (Ercanoglu et al., 2004).

2.3 Analysis

AHP has gained wide application in site selection, suitability analysis, regional planning, and landslide susceptibility analysis (Ayalew et al 2005). Using this method, each layer is broken into smaller factors, and then these factors are compared based on their importance. For comparison of importance of factors relative to each other, each factor is rated against every other factor by assigning a relative dominant value between 1 and 9 (Table 1).

Description Dominant	Value
Equal importance	1
Moderate prevalence of one over another	3
Strong or essential prevalence	5
Very strong or demonstrated prevalence	7
Extremely high prevalence	9
Intermediate values	2,4,6,8

Table 1. Fundamental scale for pair-wise comparisons (after Saaty and Vargas (2001))

In order to establish a pair-wise comparison matrix (A), factors of each level and their weights are shown as: A_1, A_2, \dots, A_n and w_1, w_2, \dots, w_n . The relative importance of a_i and a_j is shown as a_{ij} . The pair-wise comparison matrix of factors A_1, A_2, \dots, A_n as $A=[a_{ij}]$ is expressed as:

$$A = [a_{ij}]_{n \times n} = \begin{bmatrix} 1 & a_{12} & \dots & a_{1n} \\ a_{21} & 1 & \dots & a_{2n} \\ \vdots & \vdots & \ddots & \vdots \\ a_{n1} & a_{n2} & \dots & 1 \end{bmatrix} = \begin{bmatrix} 1 & \frac{w_1}{w_2} & \dots & \frac{w_1}{w_n} \\ \frac{w_2}{w_1} & 1 & \dots & \frac{w_2}{w_n} \\ \vdots & \vdots & \ddots & \vdots \\ \frac{w_n}{w_1} & \frac{w_n}{w_2} & \dots & 1 \end{bmatrix} \quad (1)$$

In this matrix, the element, $a_{ij} = 1/a_{ji}$ and thus, when $i=j$, $a_{ij}=1$. A matrix is normalized using:

$$a'_{ij} = \frac{a_{ij}}{\sum_{i=1}^n a_{ij}} \quad i, j = 1, 2, \dots, n \quad (2)$$

And finally, weights of factors are computed using

$$w_i = \left(\frac{1}{n}\right) \sum_{j=1}^n a'_{ij} \quad i = 1, 2, \dots, n \quad (3)$$

In matrix-based pair-wise comparison, if the factor on the horizontal axis is more important than the factor on the vertical axis, this value varies between 1 and 9. Conversely, the value varies between the reciprocals 1/2 and 1/9 (Table 1). In AHP, for checking consistency of matrix, consistency ratio is used, which depends on the number of parameters. The consistency ratio (CR) is obtained by comparing the consistency index (CI) with

average random consistency index (RI). The consistency ratio is defined as: $CR = CI/RI$

The consistency index of a matrix of comparisons is given by:

$$\text{Consistency Index (CI)} = \frac{\lambda_{\max} - n}{n - 1} \quad (4)$$

And the average random consistency index (R.I.) is derived from a sample of randomly generated reciprocal matrices using the scales 1/9, 1/8, ..., 8 and 9 (see Table 2).

Number of factors	RI	Number of factors	RI
1	0.00	6	1.25
2	0.00	7	1.35
3	0.52	8	1.40
4	0.89	9	1.45
5	1.11	10	1.49

Table 2: Average random consistency index (RI)

The AHP methodology consists of pairwise comparison of all possible pairs of factors and try to synthesize the judgments to determine the weights (Saaty, 2001). In this study the relative rating for the dominance between each pair of factors was guided by expert knowledge. A comparison matrix of scores was created (Table 3). In this study, the value of CR is obtained by the ratio between the values of the indexes CI (matrix's consistency index, whose expression is shown in Eq. (4)) and a random index RI, which is the average consistency index

The consistency ratio (CR) calculated is 0.053, therefore less than 0.1.

Weight-decision matrix	Precipitation	Slope	Lithology	Aspect	Land cover	Distance from road	Fracturation	Streams	Criteria Weigh
Precipitation	1	2	2	3	4	3	3	4	0,2551
Slope	0,5	1	3	3	3	4	3	4	0,2251
Lithology	0,5	0,33	1	3	2	6	4	4	0,1854
Aspect	0,33	0,33	0,33	1	2	2	1	2	0,0873
Land cover	0,25	0,33	0,5	0,5	1	2	2	2	0,0806
Distance from	0,33	0,25	0,16	0,5	0,5	1	2	1	0,0584
Fracturation	0,33	0,33	0,25	1	0,5	0,5	1	2	0,0625
Streams	0,25	0,25	0,25	0,5	0,5	1	0,5	1	0,0457

Table 3. Pairwise comparison matrix of scores for calculating weight

After obtaining weight of each parameters, they is multiplied in the map calculated by criteria weights by applying the following Eq (5):

$$S = \sum_i w_i x_i * \prod_j c_j \quad (5)$$

with S is the linear combination of factor values (x_i) weighted by weights (w_i) and multiplied by the product of constraints (c_j)

3. Results and discussion

The analysis of the susceptibility map obtained (Figure 6) shows that areas with high susceptibility to sliding are largely concentrated in the western part of the studied area at the cliffs area, due to local environmental conditions.

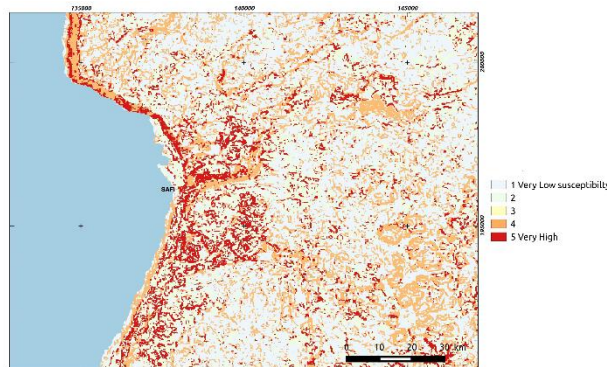


Figure 6: Landslide susceptibility maps

Landslide research and susceptibility mapping is an essential component of hazard management. Comprehending the processes of occurring landslide provides fundamental knowledge about the evolution of landscapes and decreasing the risk due to landslides. There are various methods for landslide susceptibility mapping.

In this study, AHP method was used to prepare landslide susceptibility maps in The coastal area of Safi (Morocco). This region is a landslide prone zone because of its own characteristics including the cliffs topography, climate conditions, geology and geomorphology structures. In this study area, using field records, 170 landslides have been determined. Majority of these landslides occurred because of losses of land cover, slope topographic, fracturation, absence of végétation, erosion, and unsystematic road construction in marl and clay sediment.

After preparing landslide inventory map, eight layers have been considered. To confirm the practicality of the results, the susceptibility map was compared with existing landslide. The results demonstrated that the active landslide zones had a high correlation to the high and very high susceptibility class of map. The AHP map shows 80% of the active landslide reas located at high and very high susceptibility are generally related to gray marl rocks that are very sensitive in the neighborhood of water. Based on this study, it can be stated that the high and very high susceptibility landslide zones identified by the AHP method, can predict potential landslide areas in the reality. The result of this study shows, that when field conditions are properly determined by good proficiency, the AHP method can give more truly results (Moradi et al., 2012).

4. References

Ayalew, L., Yamagishi, H., 2005. The application of GIS-based logistic regression for landslide susceptibility mapping in the Kakuda–Yahiko Mountains, Central Japan. *Geomorphology.*, 65, 15–31.

- Clerici, A., Perego, S., Tellini, C., Vescovi, P., 2002. A procedure for landslide susceptibility zonation by the conditional analysis method. *Geomorphology*, 48, 349–364.
- Dai, F.C., Lee, C.F., Li, J., Xu, Z.W., 2001. Assessment of landslide susceptibility on the natural terrain of Island, Hong Kong. *Environmental Geology*, 43 (3): 381–391.
- El Bchari F. , Theilen-Willige B., Ait Malek H. 2014. *Earth Sciences*, 3(3): 76-84
- Ercanoglu, M., Gokceoglu, C., 2004. Use of fuzzy relations to produce landslide susceptibility map of a landslide prone area (West Black Sea Region, Turkey). *Engineering Geology*, 75 (3–4): 229–250.
- Ermini L, Catani F, Casagli N (2005) Artificial neural networks applied to landslide susceptibility assessment. *Geomorphology* 66:327–343
- Fell, R., Corominas, J., Bonnard, CH., Cascini, C., Leroi, E., Z. Savage, W.,(2008a). Guidelines for landslide susceptibility, hazard and risk zoning for land use planning. *Engineering Geology*, 102, 85–98.
- Gigout M (1951) Etudes géologiques sur la Méséta marocaine occidentale (Arrière-pays de Casablanca, Mazagan et Safi). Travaux de l'Institut Scientifique Chérifien N° 3, et Notes et Mémoires du Service Géologique. N° 86. Rabat, Imprimerie Maroc-Matin, 2 tomes, 507 p
- Gigout M (1954) Carte géologique de la Meseta entre Mechrâ Benâbbou et Safi (Abda, Doukkala et massif des Rhamna) au 1/200.000, Notes et Mémoires du Service Géologique du Maroc, N° 84
- Iovine G., 2008, Mud-flow and lava-flow susceptibility and hazard mapping through numerical modelling, GIS techniques, historical and geo-environmental analyses. In: Proceedings of the iEMSs 4th biennial meeting, international congress on environmental modelling ANS software: integrating sciences and information technology for environmental assessment and decision making (iEMSs2008), vol 3, pp 1447–1460
- Irigaray C (1995) Movimientos de Ladera: Inventario, Análisis y Cartografía de la Susceptibilidad Mediante un Sistema de Información Geográfica. Aplicación a las Zonas de Colmenar (Málaga), Rute (Córdoba) y Montefrío (Granada). Unpublished PhD Thesis. Department of Civil Engineering. University of Granada, Spain
- Irigaray C, Fernández T, El Hamdouni R, Chacón J (1999) Verification of landslide susceptibility mapping. A case study. *Earth Surf Process Landf* 24:537–544
- Irigaray C, Fernández T, El Hamdouni R, Chacón J (2007) Evaluation and validation of landslide-susceptibility maps obtained by a GIS matrix method: examples from the Betic Cordillera (southern Spain). *Nat Hazards* 41:61–79
- Jiménez-Perálvarez JD, Irigaray C, El Hamdouni R, Chacón J (2009) Building models for automatic landslide-susceptibility analysis, mapping and validation in ArcGIS. *Nat Hazards* 50:571–590. <https://doi.org/10.1007/s11069-008-9305-8>
- Jiménez-Perálvarez JD (2012) Movimientos de ladera en la vertiente meridional de Sierra Nevada (Granada, España): Identificación, análisis y cartografía de susceptibilidad y peligrosidad mediante SIG, Tesis Doctoral, Programa de doctorado : Ciencias y Tecnología del Medio Ambiente, Universidad de Granada, E.T.S. de Ingenieros de Caminos, Canales y Puertos, Grupo de Investigaciones Medioambientales: Riesgos Geológicos e Ingeniería del Terreno; RNM-121, Departamento de Ingeniería Civil, 210 p
- Komac, M. 2006. A landslide susceptibility model using the Analytical Hierarchy Process method and multivariate statistics in perialpine Slovenia. *Geomorphology*, 74, 17–28.
- Lee, S., Choi, J., Min, K., 2004a. Probabilistic landslide hazard mapping using GIS and remote sensing data at Boun, Korea. *International Journal of Remote Sensing*, 25 (11): 2037–2052.
- Lee, S., 2005. Application of logistic regression model and its validation for landslide susceptibility mapping using GIS and remote sensing data. *International Journal of Remote Sensing*, 26 (7): 1477–1491.
- Malczewski, J., 1999. GIS and Multi-criteria Decision Analysis, JohnWiley and Sons, New York.
- Mittal, S.K.,Singh, M., Kapur, P.,Sharma, B.K., Shamshi, M.A., 2008. Design nd development of instrument network for landslide monitoring, an issue an early warning. *Journal of Scientific &Industrial research* 67,361–365.
- Moradi M., Bazyar M. H., Zargham M. 2012 *Journal of Basic and Applied Scientific Research*, 2(7), 6715-6723.
- Nandi, A., Shakoor, A. 2009. A GIS-based landslide susceptibility evaluation using bivariate and multivariate statistical analyses. *Engineering Geology*, 110 (1-2): 11-20.
- Nielsen TH, Wrigth RH, Vlastic TC, Spangle WE., 1979, Relative slope stability and land-use planning in the San Francisco Bay region, California. US geological survey professional paper 944
- Ouadia M (1998) Les formations plio-quaternaires dans le domaine mesetien occidental du Maroc entre Casablanca et Safi: géomorphologie, sédimentologie, paléoenvironnements quaternaires et évolution actuelle. Thèse de Doctorat d'Etat Ès Sciences, Université Mohamed V, N° 1646, Rabat, 321 p.
- Pachauri, A.K., Pant, M., 1992. Landslide hazard mapping based on geological attributes. *Engineering Geology*, 32, 81–100.
- Regmi NR, Giardino JR, Vitek JD (2010) Assessing susceptibility to landslides: using models to understand

observed changes in slopes. *Geomorphology* 122(1–2):25–38.

A case study on natural slopes of Hong Kong, China. *Geomorphology*, 101, 572–582.

Regmi et al. 2010; Ruff and Rohn 2008, Canani et al. (2008). modeling approaches (Perriello Zampelli et al. 2012), Fuzzy Analytical Hierarchy Process (FAHP) (Shadman Roodposhti et al., 2010), Geographically weighted principal component analysis (Faraji Sabokbar et al., 2014),

Roch E (1930) Carte géologique provisoire des Abda et des Djebilet occidentales, au 1/200000. Notes et mémoires du Service Mines et Carte géologique. Maroc, n°10

Roch E (1950) Histoire stratigraphique du Maroc. Notes et mémoires du Service Géologique n° 80. Les Frères DOULADOURE. Toulouse, 435 p

Ruff M, Rohn J (2008) Susceptibility analysis for slides and rockfall: an example from the Northern Calcareous Alps (Vorarlberg, Austria). *Environ Geol* 55(2):441–452

Saaidi E (1988) Géologie du Quaternaire marocain. Société Marocaine des Editeurs Réunis, Rabat 440 p

Saaty, T.L., 1980. *The Analytical Hierarchy Process*, McGraw Hill, New York

Saaty, T.L., Vargas, G.L., 2001. *Models, Methods, Concepts, and Applications of the Analytic Hierarchy Process*, Kluwer Academic Publisher, Boston.

Saha, A.K., Gupta, R.P., Arora, M.K., 2002. GIS-based landslide hazard zonation in the Bhagirathi (Ganga) valley, Himalayas. *International Journal of Remote Sensing*, 23 (2): 357–369.

Shadman Roodposhti.M, Rahimi. S, Jafar Beglou. M., 2012, PROMETHEE II and fuzzy AHP: an enhanced GIS-based landslide susceptibility mapping, *Nat Hazards* DOI 10.1007/s11069-012-0523-8

Theilen-Willige B, El Bchari F, Malek HA, Chaibi M, Charif A, Nakhcha C, Ait Ougougdal M, Ridaoui M, Boumaggard E 2014a Remote sensing and GIS contribution to the detection of areas susceptible to natural hazards in the Safi Area, W-Morocco. In *Information and Communication Technologies for Disaster Management (ICT-DM)*, 2014 1st International Conference on (pp. 1-5). IEEE

Vargas, L.G., 1990. An overview of the analytic hierarchy process and its applications. *European Journal of Operational Research*, 48, 2–8.

Witam O 1988 Etude stratigraphique et sédimentologique de la série mésozoïque du bassin de Safi, Mémoire n°3 Thèse de Doctorat de 3ème Cycle Faculté des Sciences Marrakech. 194 p

Yalcin, A. 2008. GIS-based landslide susceptibility mapping using analytical hierarchy process and bivariate statistics in Ardesen (Turkey). *Catena*, 72, 1–12.

Yao, X., Tham, L.G., Dai, F.C., 2008. Landslide susceptibility mapping based on Support Vector Machine: